

Coupled Ocean-Wave Model Team (Team 8) Report

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HFIP Meeting, Miami, FL., 8-9 Nov. 2011

Team 8 Outline

1. Operational development at NCEP
2. COAMPS-TC coupling (NRL-MRY)
3. NOAA/ESRL sea spray flux parameterizations (ESRL and URI)
4. Sensitivity of SST cooling to ocean model resolution (URI)
5. Sensitivity of HWRF intensity forecasts to ocean coupling (AOML/PhOD and AOML/HRD)

Operational Development at NCEP

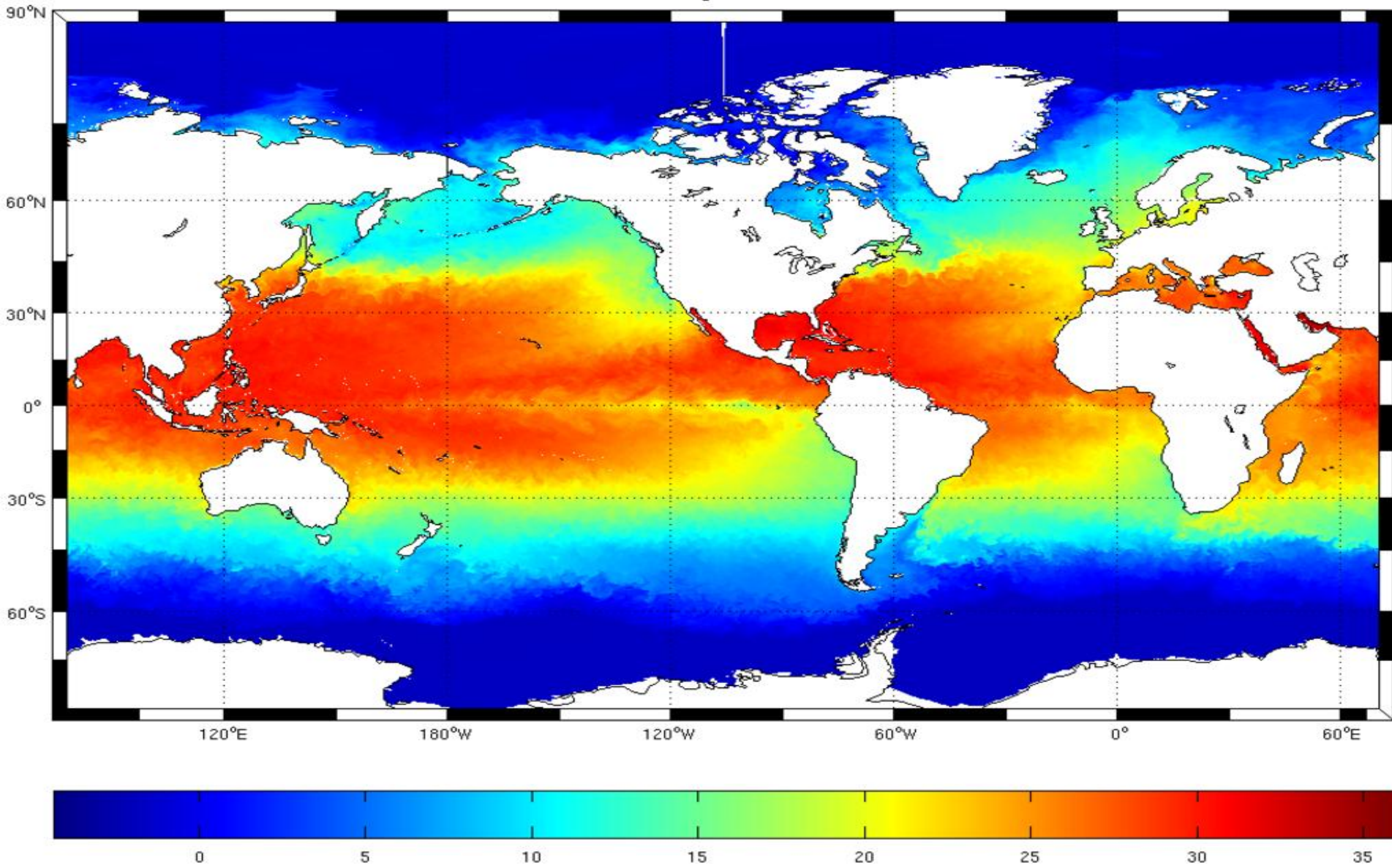
1. RTOFS-Global HYCOM transitioned into operations
2. Coupler design consideration for HYCOM-HWRF (HyHWRF)
3. Real time parallel results of HyHWRF

RTOFS-Global

- Global 1/12 degree HYCOM model implemented operationally 10/26/2011.
 - Partnership with U. S. Navy. NCODA initialization provided daily by Navy.
 - Application for hurricane modeling (HFIP)
 - Simplified ocean model initialization anywhere in the world.
 - Ocean heat content products

1/12 Degree Global Domain

Global Temperature (deg C) 20100823 n00 layer 1
NCEP/EMC/MMAB 24-Aug-2010 max: 35.96 min: -4.38



Coupler considerations

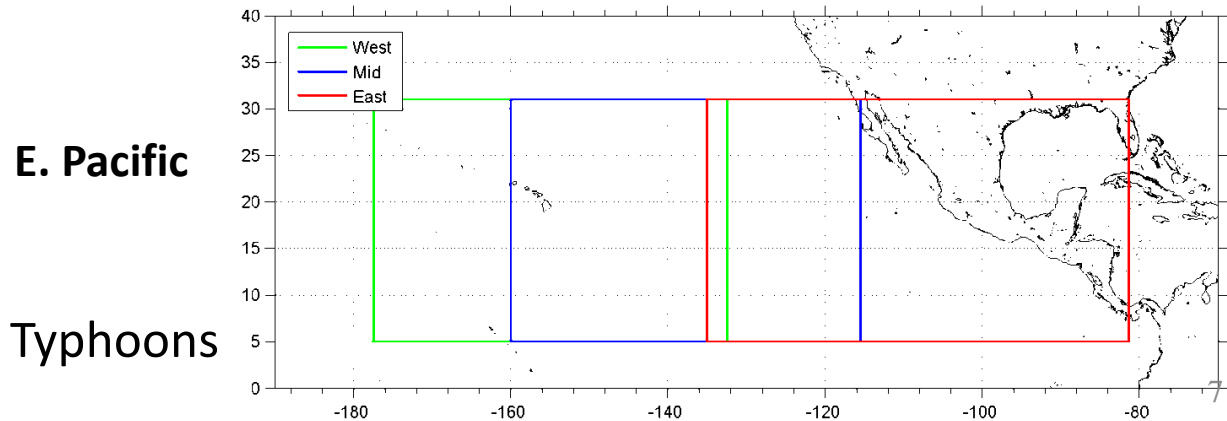
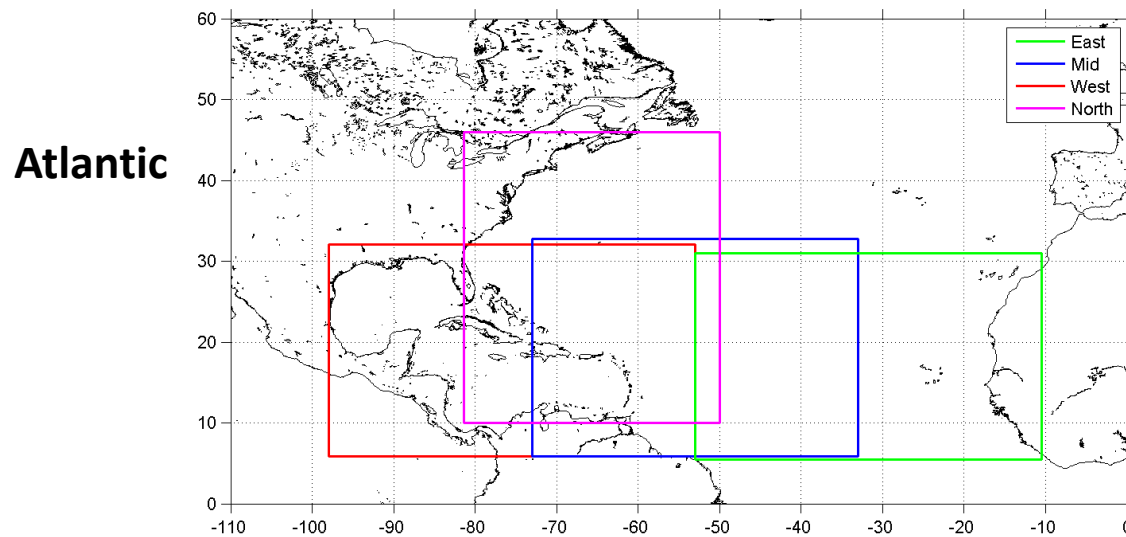
- Simple view of coupling:
 - Couple the models and improvement immediately realized
- Reality:
 - Three HWRF configurations require different optimization
 - Uncoupled (no ocean response, e.g. E.Pac) .
 - Coupled to POM (weak ocean response, e.g. Atlantic).
 - Coupled to HYCOM (strong ocean response)
 - GFS forcing used for global HYCOM is drastically different from surface fluxes produced by HyHWRF
 - Causes rapid upper-ocean T drift at start of forecasts
 - Correcting this problem requires extensive modification and tuning of WRF and HYCOM

1. New HyHWRF using 1/12-degree Global HYCOM

1-2

Accomplished/In-Progress, including

- A. Hurricane subdomain configurations in Atlantic and Pacific Basins
- B. Data Assimilation used in HyHWRF to improve initialization provided by global RTOFS HYCOM (adapt 3D VAR method from global RTOFS)
- C. Effort to utilize ESMF capability in HYCOM for future coupler



Also Western Pacific for Typhoons

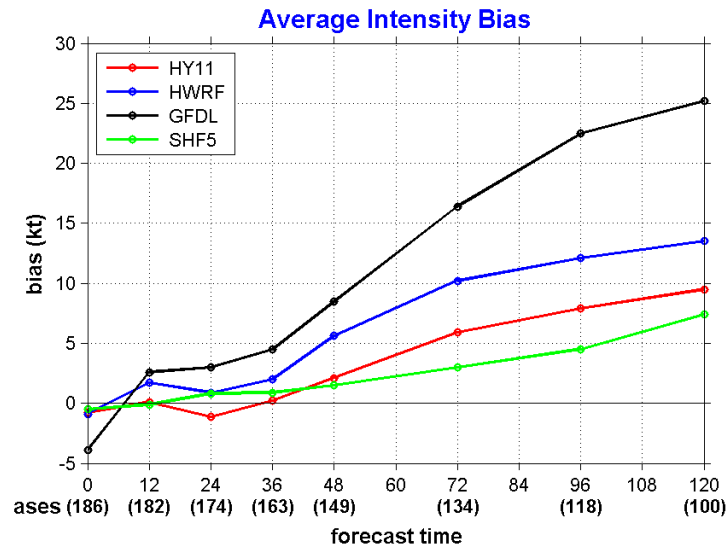
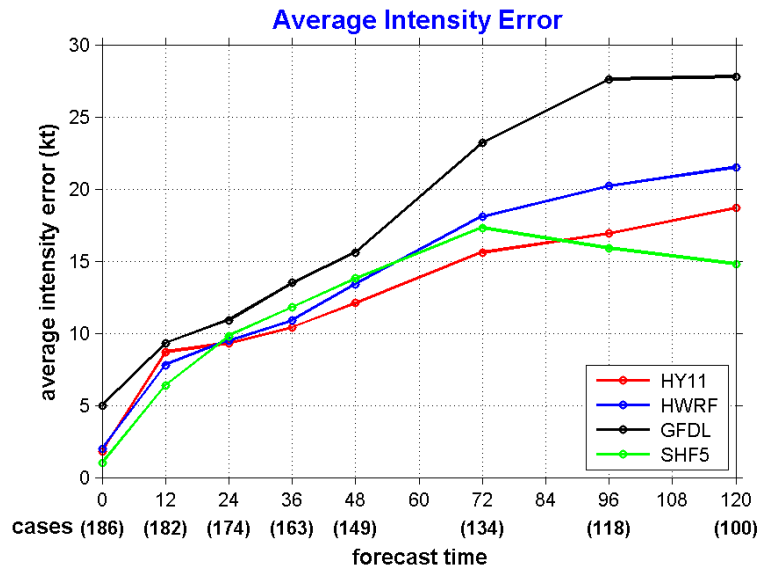
2011 Real-Time Tests of HyHWRF

➤ Summary of run results:

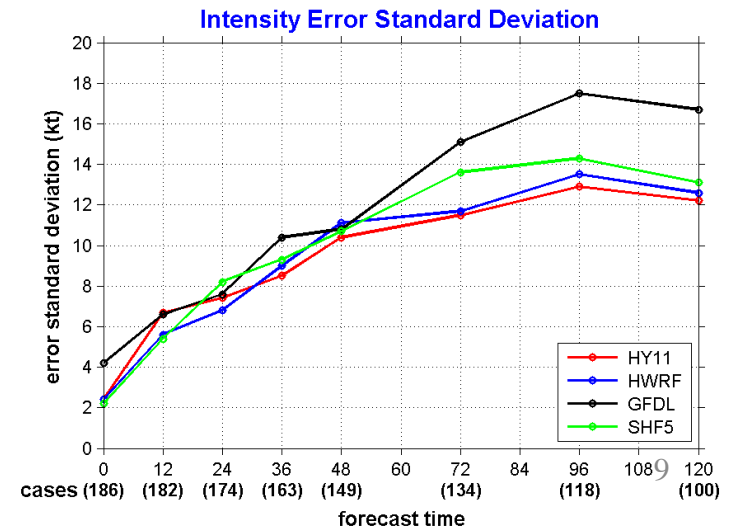
- For 6 tropical cyclones (186 cases)
 - HyHWRF shows improvement in intensity error and bias ($< \sim 5$ kt)
 - HyHWRF track forecast is highly comparable to HWRF.

Intensity Forecast for 6 TCs (186 cases): Gert07L, Irene09L, Katia12L, Maria14L, Ophelia16L, and Philippe17L

HY11=HyHWRF2011; HWRF=operational HWRF

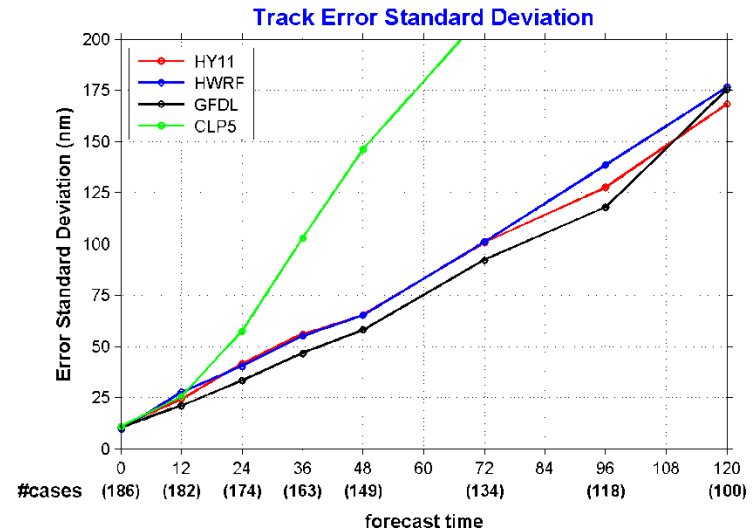
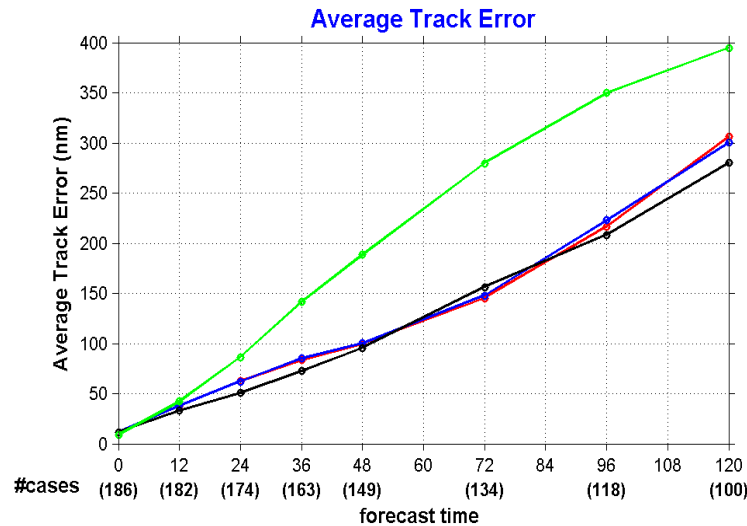


- HyHWRF average intensity error and bias are the best among participant models, except degradation at 12h in average error and negative bias at 24h.
- HyHWRF standard deviation is consistently the smallest, except 12 and 24 h.

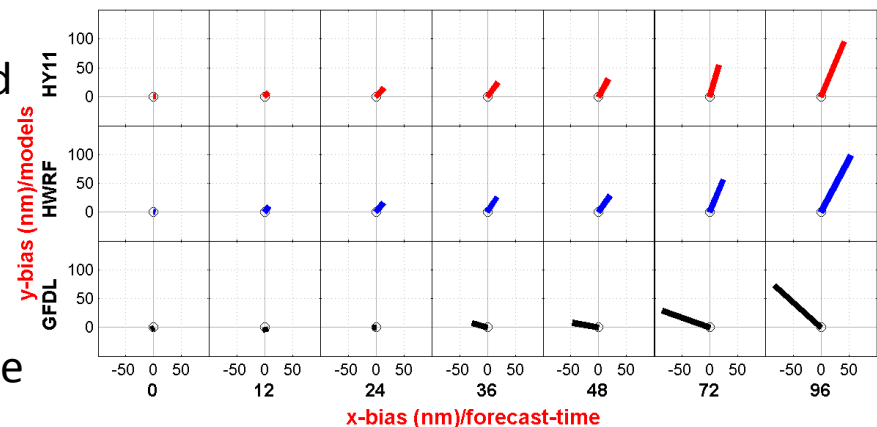


Track Forecast for 6 TCs (186 cases): Gert07L, Irene09L, Katia12L, Maria14L, Ophelia16L, and Philippe17L

HY11=HyHWRF2011; HWRF=operational HWRF



- HyHWRF exhibits comparable performance with HWRF for average track error and standard deviation (STD), except STD outperformance at 96 -120 h. GFDL shows mixed comparison but the best STD.
- HyHWRF track bias is the same northeastward as HWRF, but the bias magnitude is better than the HWRF. GFDL shows west/northwestward bias.





Coupled COAMPS-TC at NRL-MRY

Progress

- Merged new atmospheric physics including microphysics, cumulus, PBL, and sea spray from uncoupled COAMPS-TC

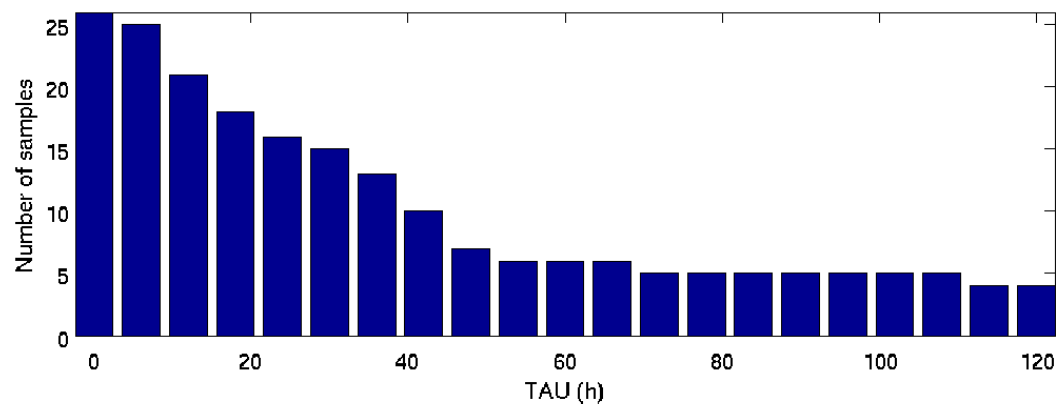
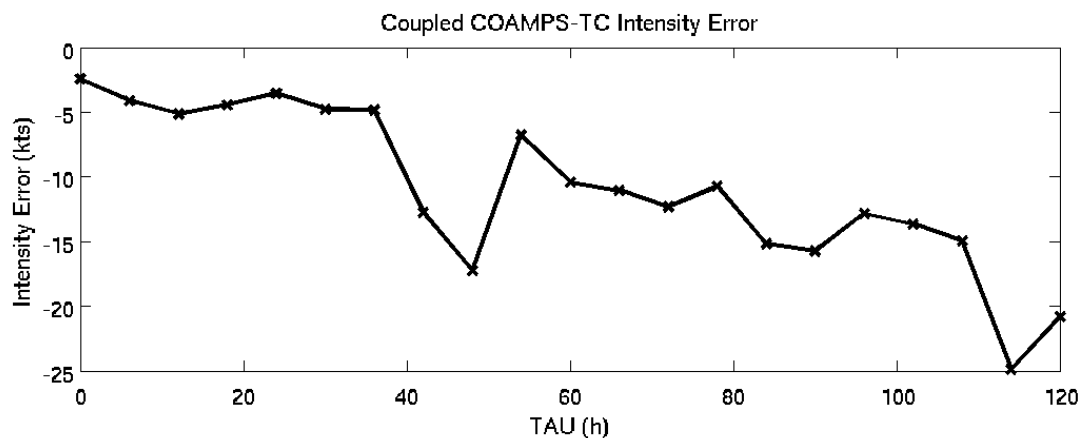
- • Performed near-real time air-ocean coupling tests with the updated new atmospheric physics in the Atlantic basin

- • Performed XBT assimilation impact studies

- • Performed air-ocean-wave coupled tests on selected tropical cyclones



Coupled COAMPS-TC Homogenous Intensity Error (09L, 12L, 14L, 16L, 17L)



Coupled COAMPS-TC has an averaged (27 samples) negative intensity bias, suggesting:

- Higher horizontal resolution may be needed for the coupled COAMPS-TC

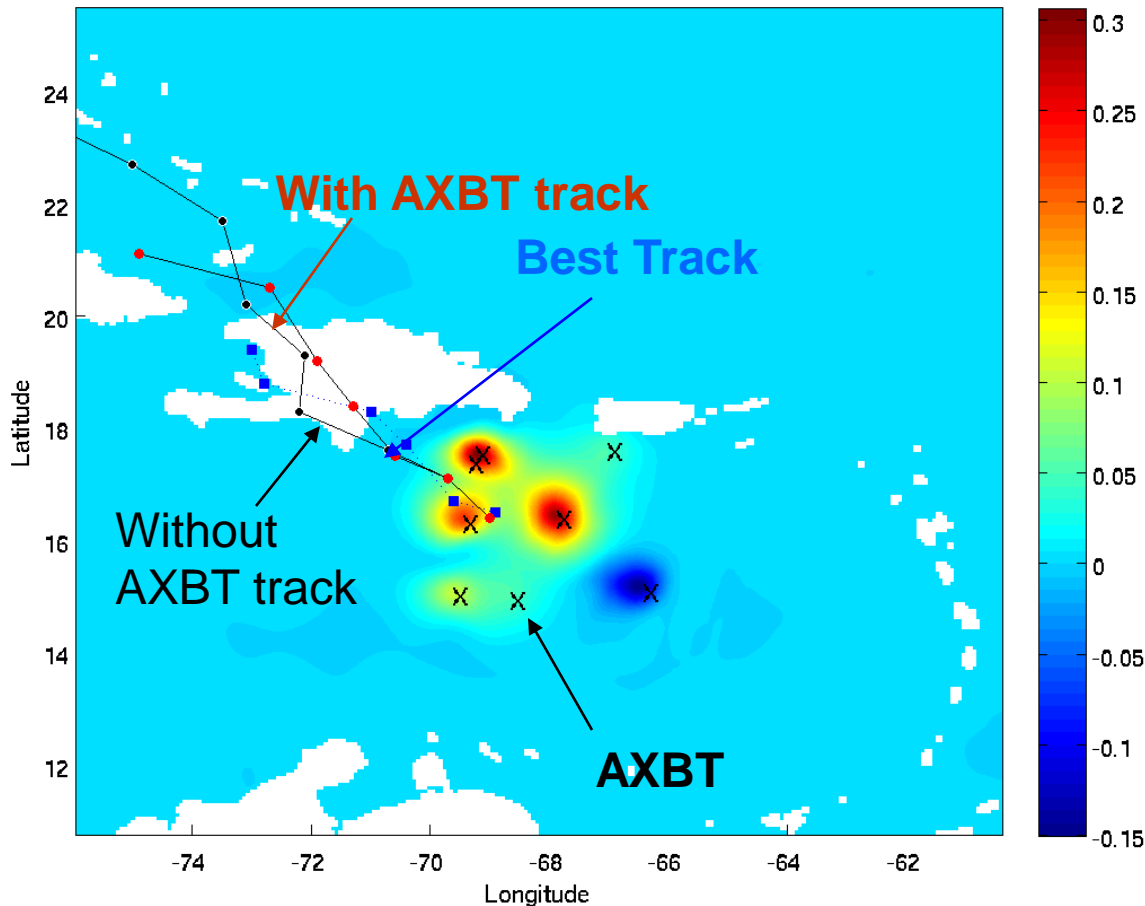
or

- Further adjustment of new atmospheric physics for 5 km coupled COAMPS-TC is needed



AXBT Impact Study – Hurricane Emily

Sea Surface Temperature Difference(C), 2011080312

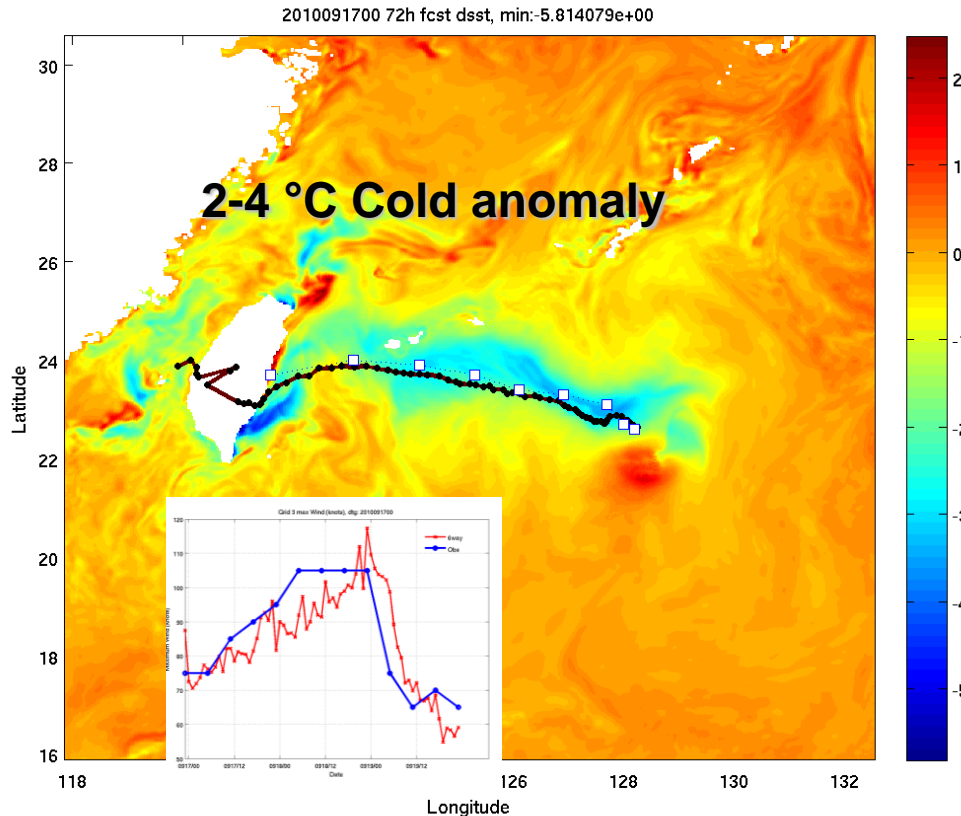


- **AXBT assimilation run has a warmer SST than the run without the AXBT**
- **Slight track difference after 12h**
- **AXBT assimilation run has a landfall location close to the observation**



Air-Ocean-Wave Coupled COAMPS-TC

High-Resolution Coupled COAMPS Simulations of Typhoon Fanapi (2010)



**Atmosphere: 27, 9,
and 3 km**

Ocean: 9 and 3 km

Wave: 1/6 degree

**Model spin-up from
2010090800**

12 h update cycle

- COAMPS forecast of Fanapi intensity is promising
- Model is 6 hr too slow due to adjustment to the unbalanced TC bogus vortex in the first 12 h

ESRL Air-Sea Coupled Modeling

in collaboration with URI

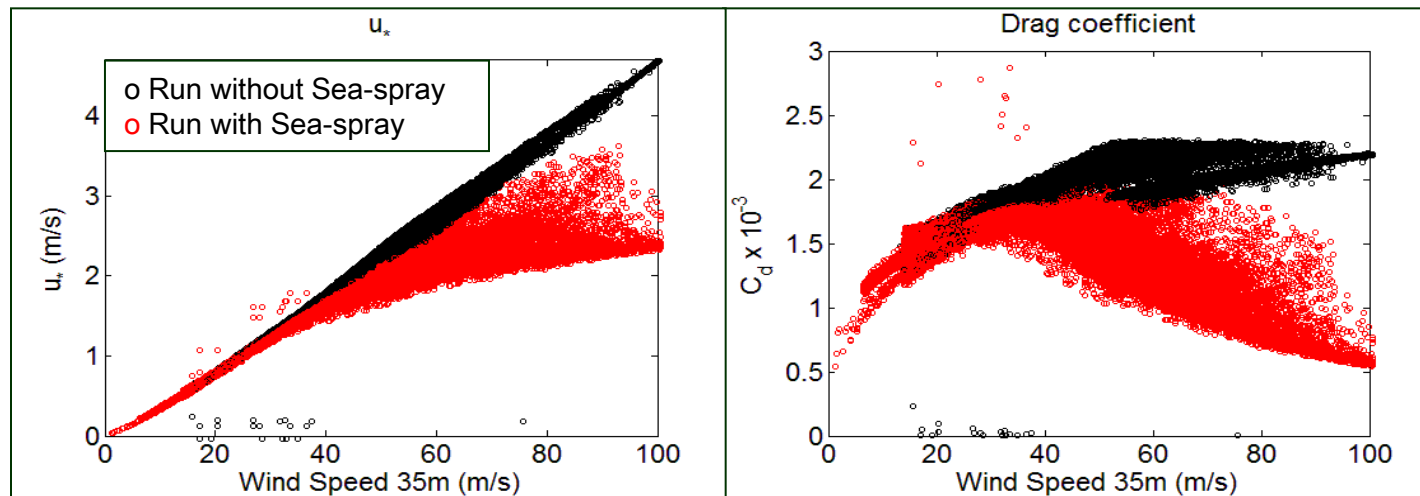
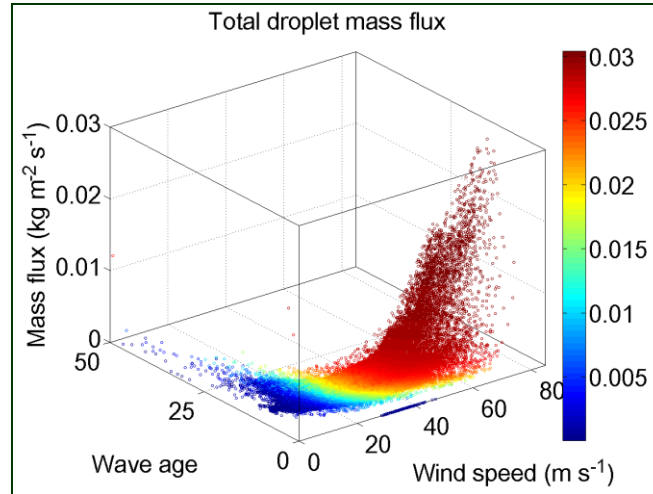
Task:

In collaboration with the URI group, we have implemented and tested the NOAA/ESRL sea spray parameterization scheme in the GFDL coupled atmosphere-wave-ocean hurricane model for an idealized case. The scheme predicts that the overall impact of sea-spray droplets on the mean winds depends on the wind speed at the level of sea-spray generation.

Results:

As the wind speed increases, the droplet size increases and the overall wind speed in the surface layer above the level of sea-spray generation increases, indicating that **the increase of droplet size due to the increase of wind speed enhances the vertical mixing**. This is consistent with observations and results from previous numerical model simulations of the microphysical characteristics of sea spray in the atmospheric boundary layer.

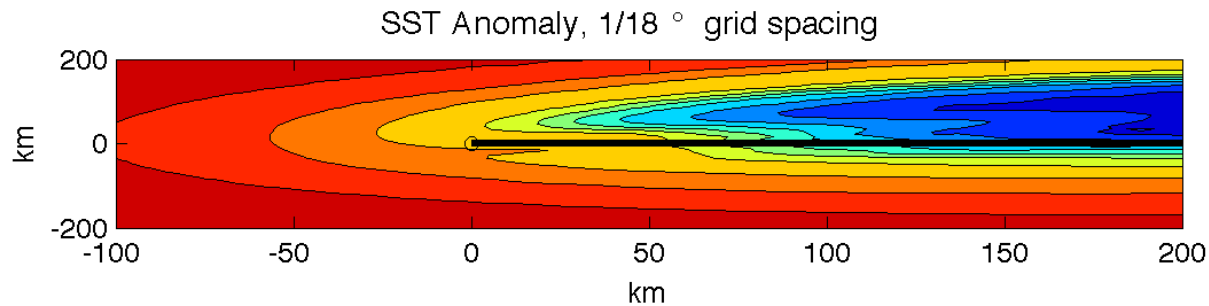
Testing the ESRL Sea-Spray Scheme in the Coupled GFDL Hurricane Model



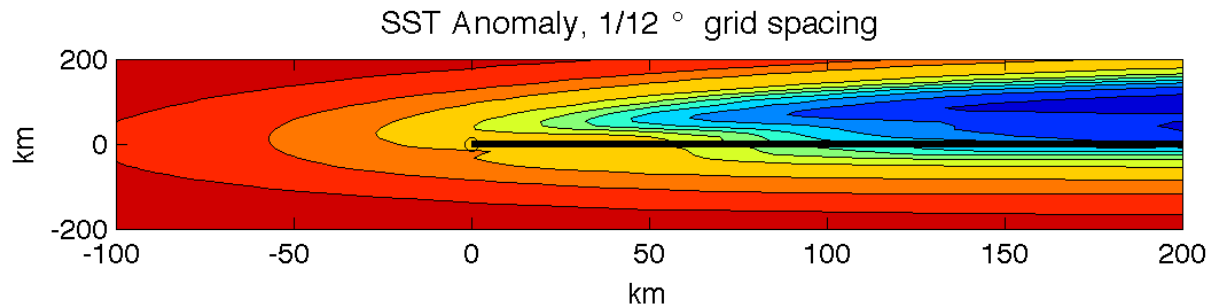
What is the optimal horizontal resolution for the ocean model?

(Idealized experiments with the Princeton Ocean Model)

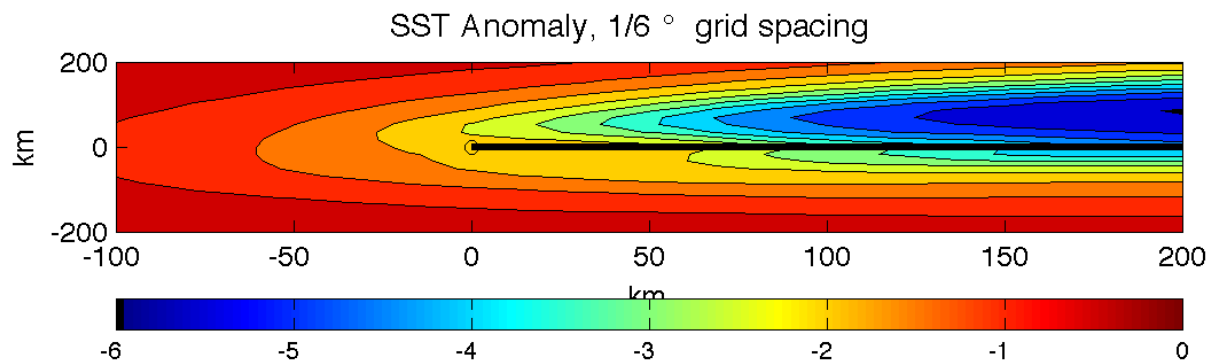
$1/18^\circ$



$1/12^\circ$



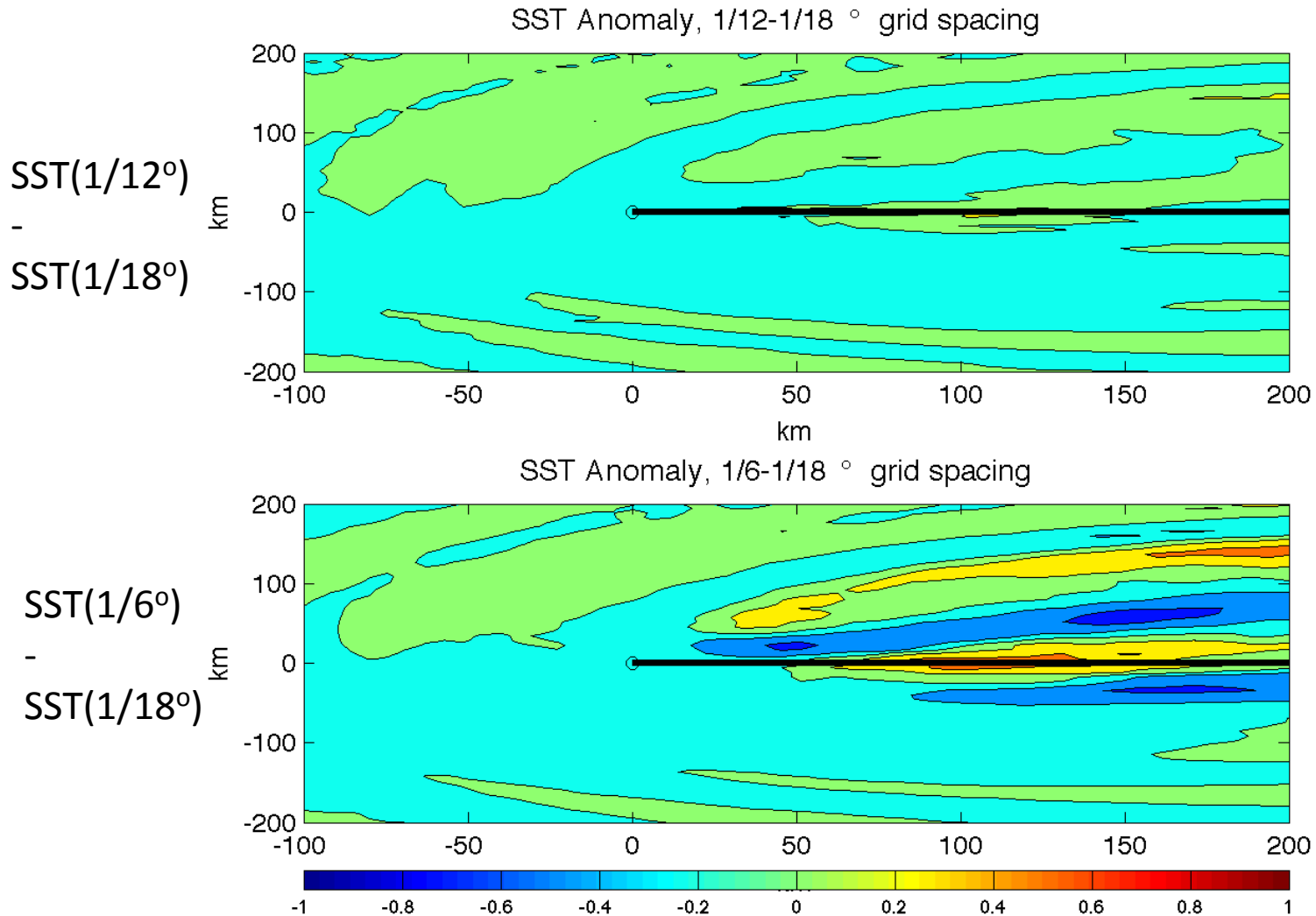
$1/6^\circ$



Hurricane:
 $R_m=55$ km,
 $V_m=55$ m/s
 $U_T=5$ m/s

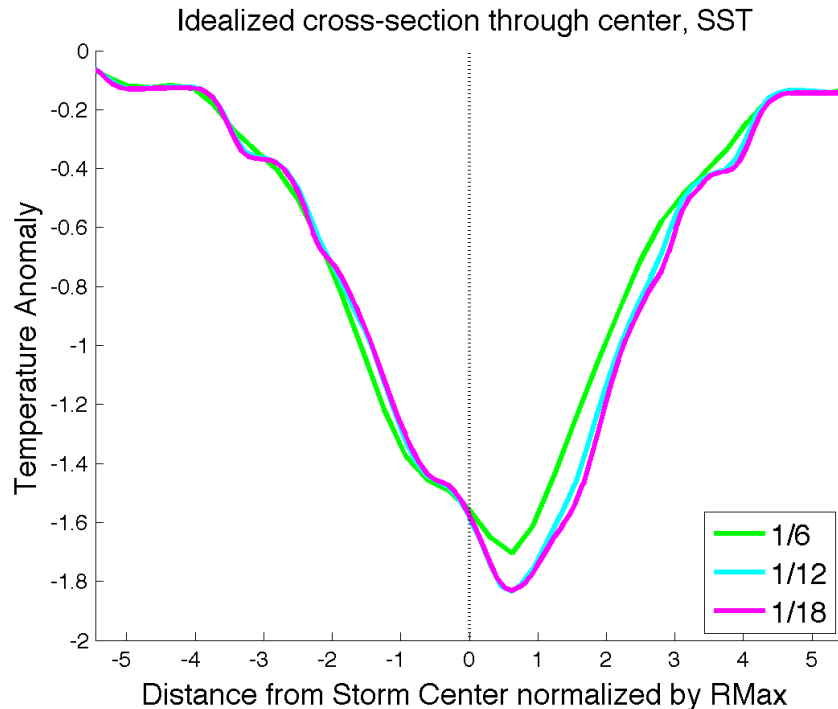
SST Differences

Between $1/6^\circ$, $1/12^\circ$, $1/18^\circ$ Resolutions

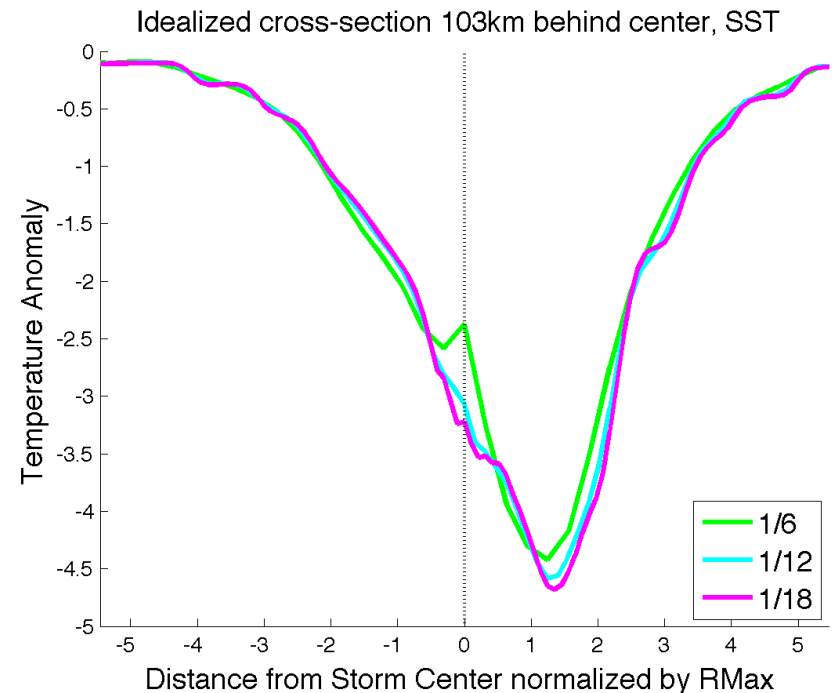


SST Cross-track Profiles

Through center



103 km behind center



Conclusions:

1. SST cooling within hurricane core in the 1/12° and 1/18° runs is larger (but not much) than in 1/6° run
2. The 1/12° and 1/18° results are very similar

Sensitivity of HWRF V3.2 to the Ocean

- Problem: poor quantitative understanding of the sensitivity of forecast intensity to changing ocean conditions in coupled forecast models (when/where is the ocean important?)
- Questions:
 - What is the dependence of quasi-equilibrium storm intensity on ocean heat content?
 - How is forecast intensity affected by changing ocean heat content and SST cooling rate?
 - Quantify the magnitude and time scale of the intensity response
 - How do storm parameters affect these results?
 - Storm size and translation speed
- Approach: Perform idealized HWRF V3.2 study minimizing impact of atmospheric processes that affect intensity
 - Magnitude and pattern of SST cooling is the dominant large-scale process affecting intensity

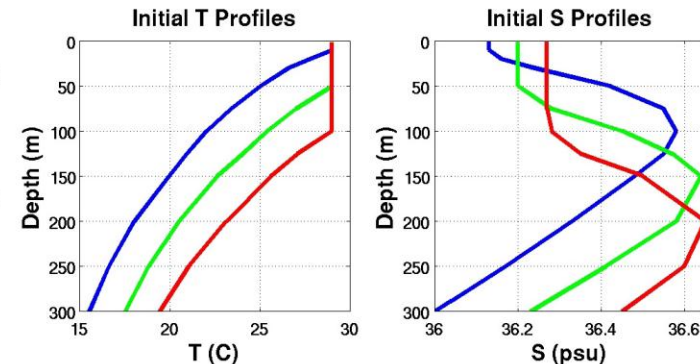
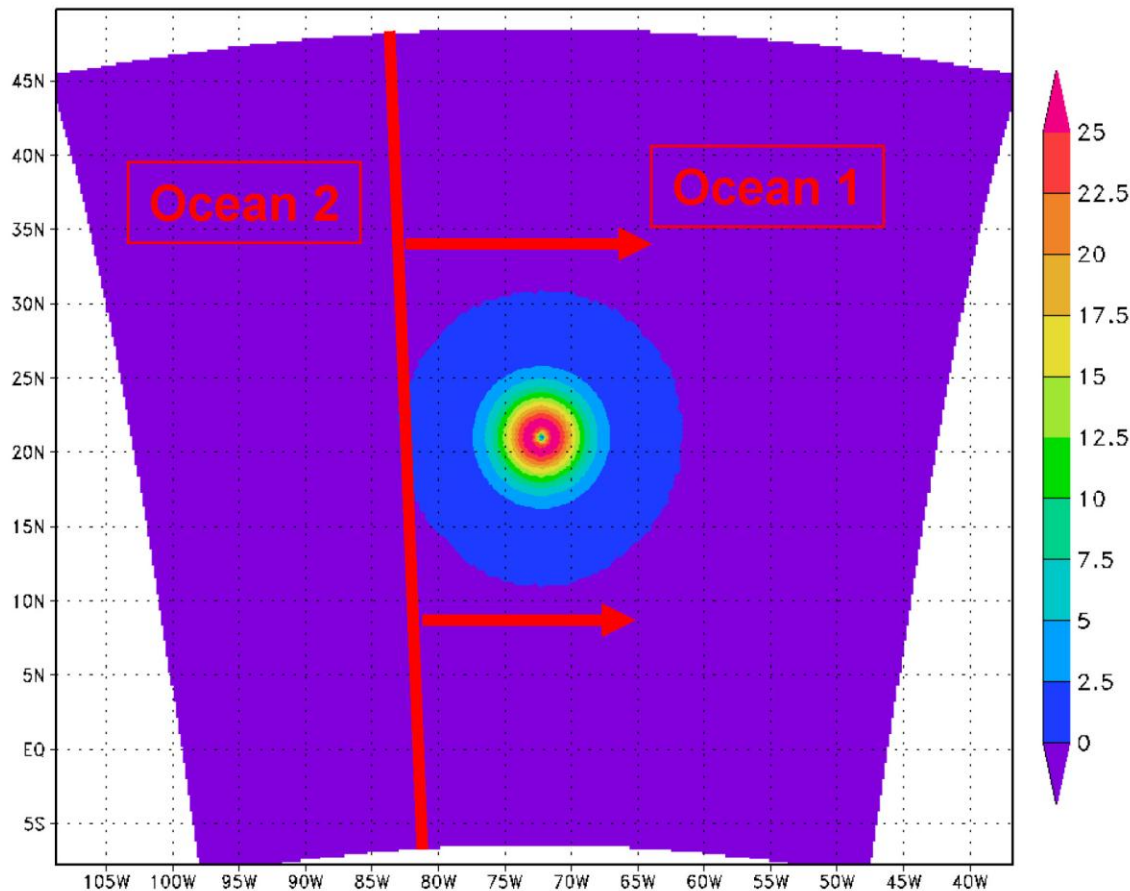
Idealized Study

- NOAA/HRD HWRF Version 3.2, 27-9-3 resolution
- Idealized atmosphere
 - Background is horizontally uniform and at rest
 - Initialized with a weak axisymmetric vortex highly favorable for intensification
 - Resulting storm remains approximately stationary
- Idealized Ocean
 - Embed one-dimensional mixing models extracted from HYCOM into HWRF 3.2
 - Initialized using specified ocean profiles
 - Ocean is bodily advected eastward past the storm to mimic storm translation
 - Use of uniform Trade Winds to control storm speed failed

Ocean Model Initialization

Initialized with either a homogeneous ocean (ocean 1 = ocean 2) or a dual ocean (ocean 1 \neq ocean 2)

Initialized Wind Speed (m/s)



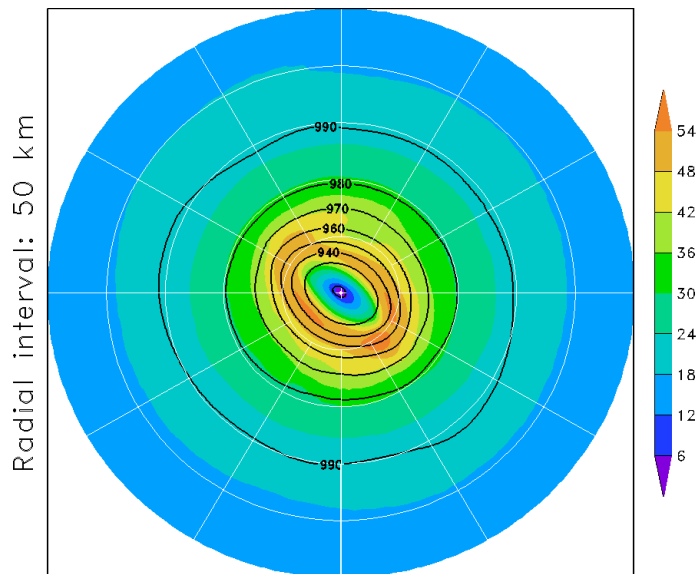
Initial Profiles:

Cool Ocean (OHC=25)

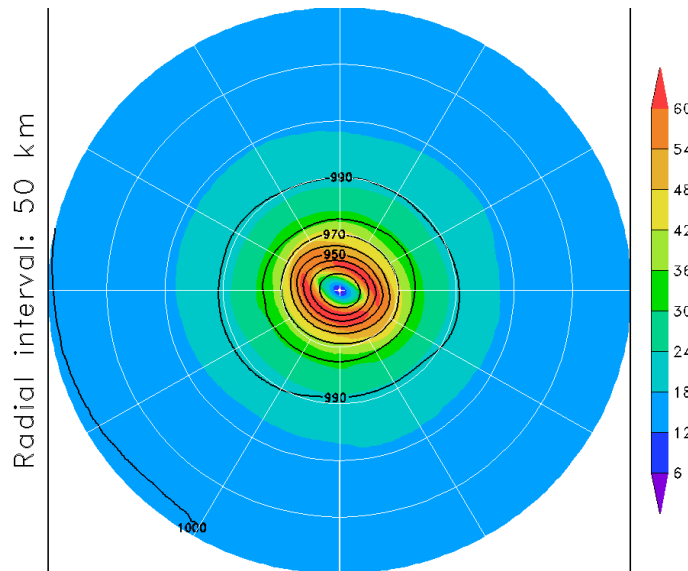
Warm Ocean (OHC=85)

Hot Ocean (OHC=148)

Large Storm



Medium Storm

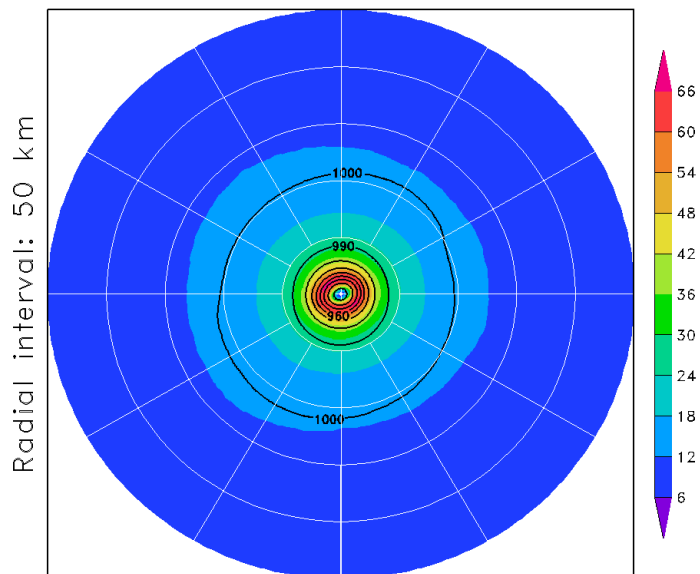


**Uncoupled
3 storm sizes**

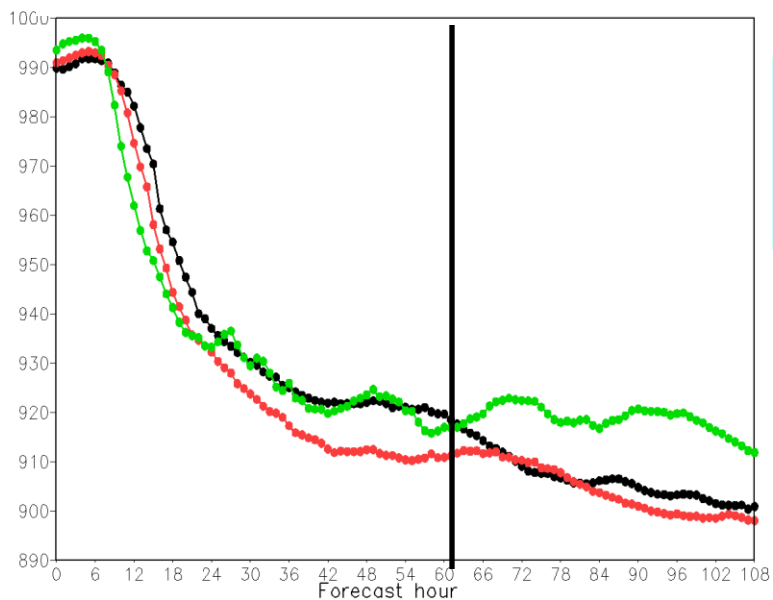
(forecast hour 60)

**Wind speed (color)
MSLP (contours)**

Small Storm

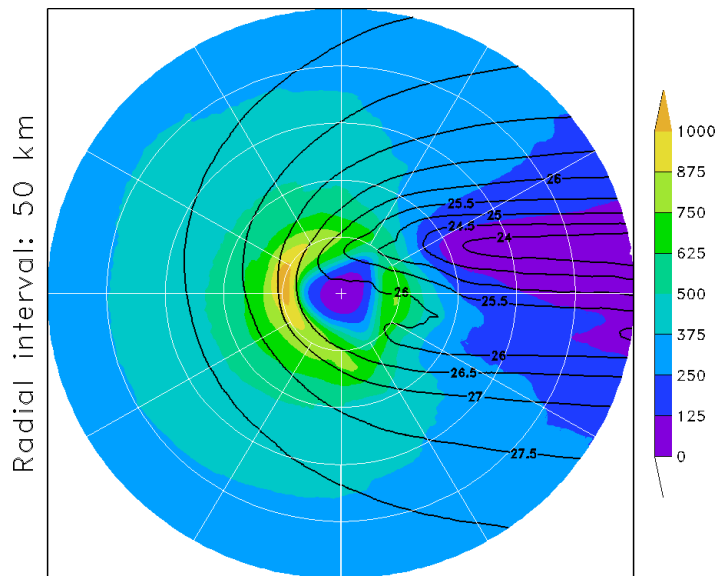


Minimum MSLP

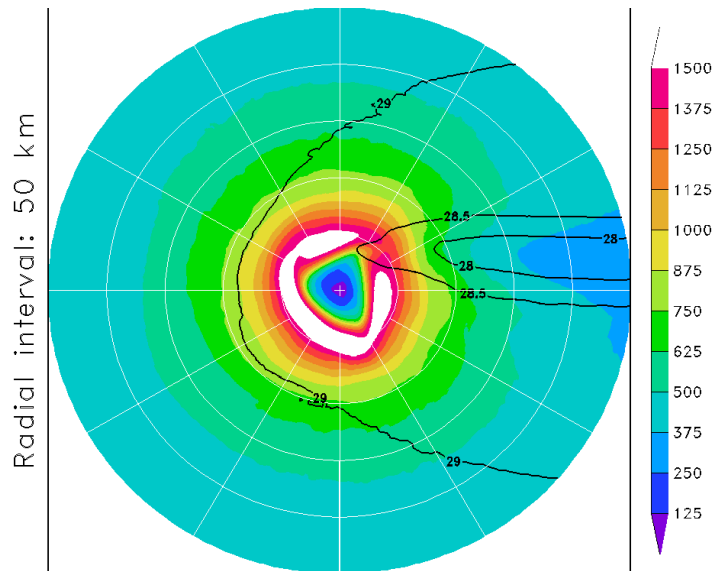


**Large storm
Medium storm
Small storm**

Cool Ocean



Hot Ocean



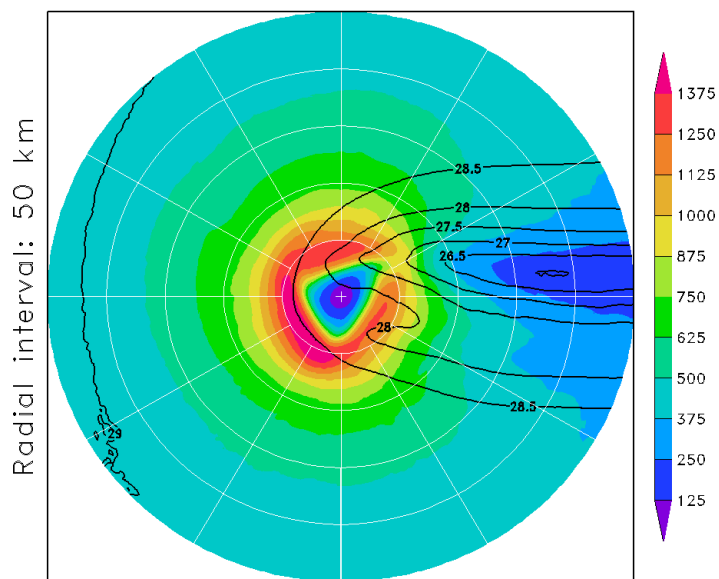
Ocean Heat Content

(forecast hour 60)

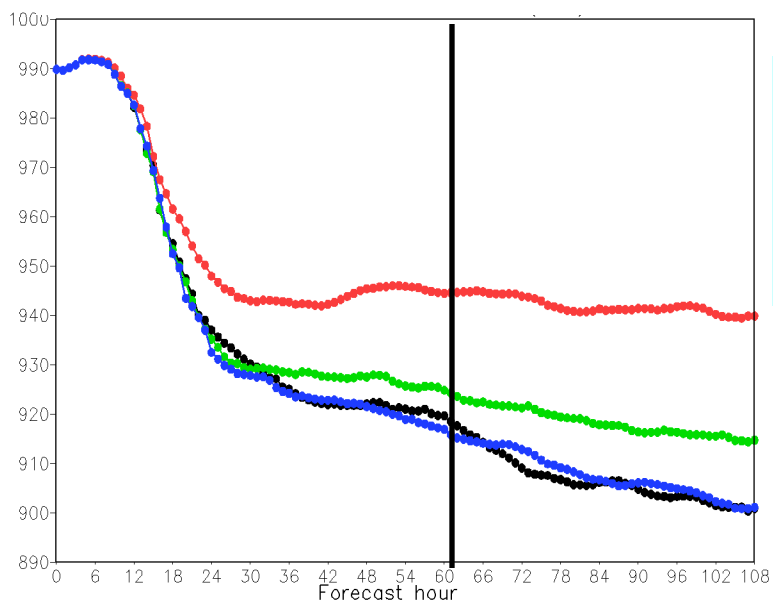
Enthalpy flux (color)

SST (contours)

Warm Ocean



Minimum MSLP



Uncoupled

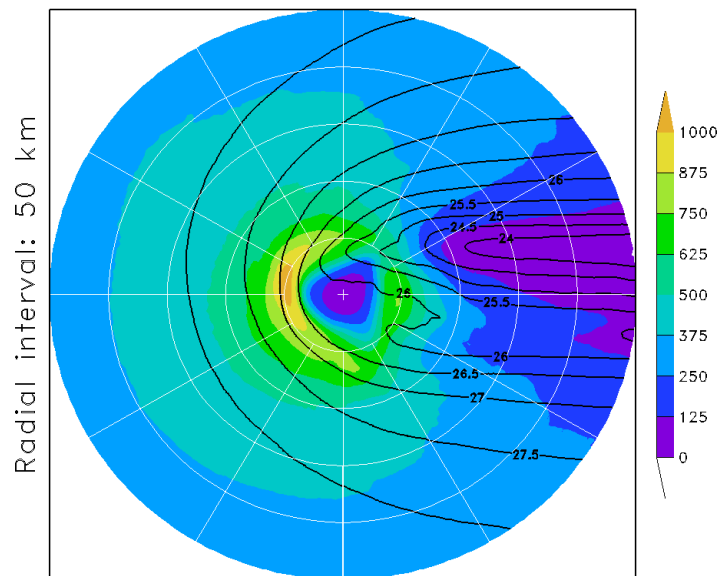
Cool ocean

Warm ocean

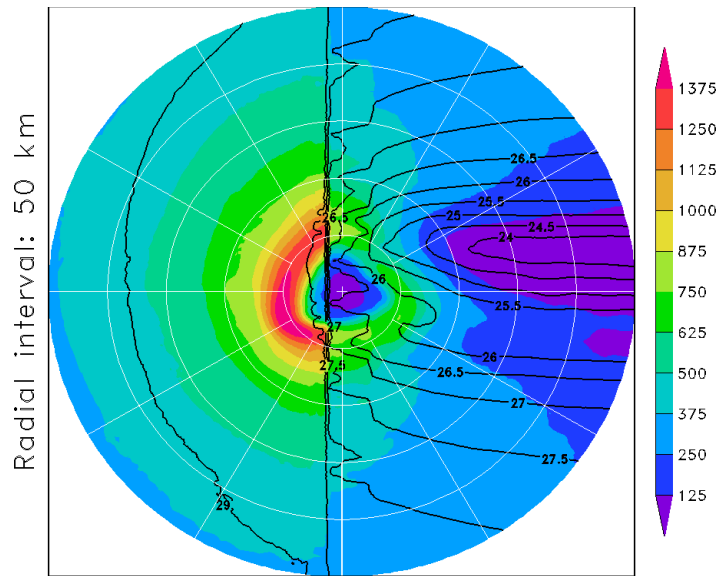
Hot ocean

**Large storm
4 m/s speed**

Cool Ocean



Cool → Warm Ocean

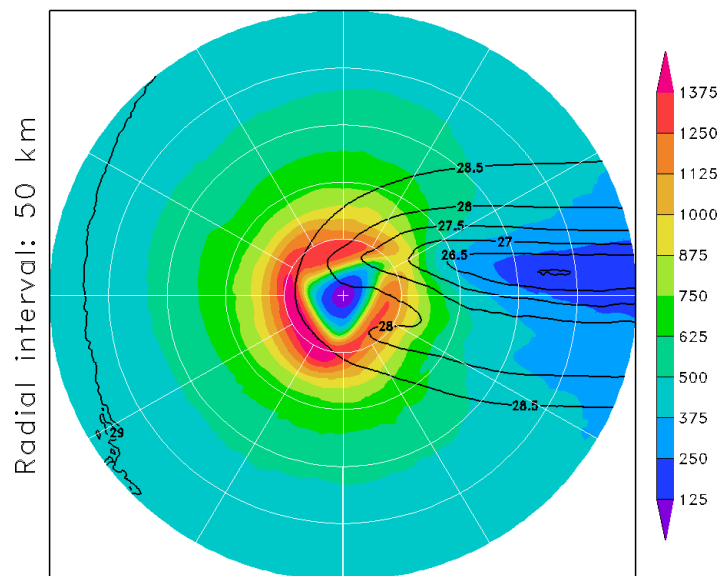


**Changing
Ocean Heat
Content**

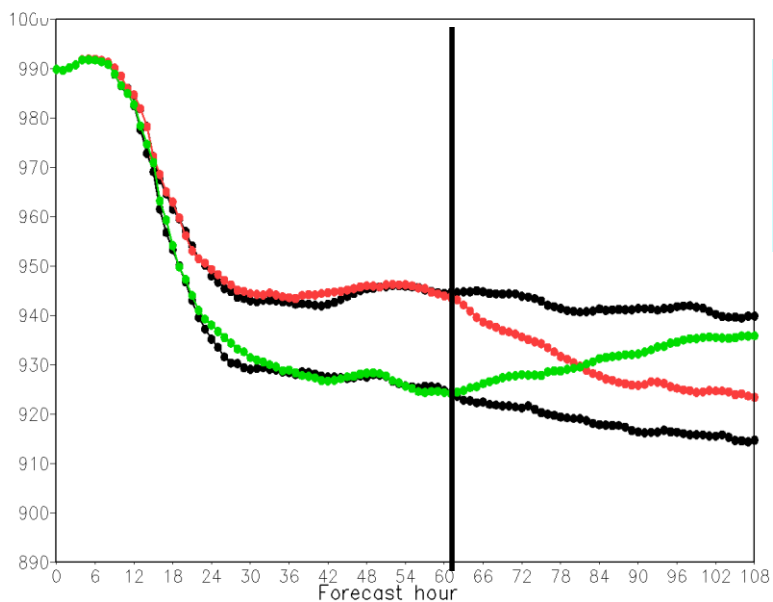
(forecast hour 60)

**Enthalpy flux
(color)
SST (contours)**

Warm Ocean



Minimum MSLP



**Large storm
Medium storm
Small storm**

**Large storm
4 m/s speed**

Limitations of 1-D Ocean Models

- Significant limitations of 1-D models, even in a horizontally uniform ocean
 - Does not properly reproduce near-inertial wave response
 - Local pure inertial current response produced by 1-D model remains trapped in the OML, overestimating shear and entrainment
 - Does not account for wind-driven upwelling
 - Very important for slow and near-stationary storms (e.g. Yablonsky and Ginis, 2009)
 - Also running idealized experiments where $w(z)$ is diagnosed from continuity eq. at each grid point and used to estimate upwelling contribution to SST change

Final Comments

- Idealized study described here will be extended to Hy-HWRF (3-D ocean)
- Similar idealized studies should be performed with other coupled forecast models
- Realistic ocean coupling studies need to be performed as soon as possible (all models)
 - Thorough scientific analysis of how ocean coupling (changes in the pattern and rate of SST cooling and enthalpy flux) affects storm intensity
 - Need high-quality upper-ocean observations beneath storms to evaluate the SST cooling patterns predicted by forecast models